

Minimum age of the Neoproterozoic Seven Hundred Acre Island Formation and the tectonic setting of the Islesboro Formation, Islesboro block, Maine

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Two platformal stratigraphic sequences occur on Islesboro, Penobscot Bay, Maine. The older Seven Hundred Acre Island Formation is at least 200 m thick, its base is not exposed, and it makes up fault-bounded blocks of siliceous colour-banded dolomitic marble, muscovite-rich quartzite, coarse-grained splendent muscovite-garnet-stauroilite-andalusite schist, and calcareous metapelite, with minor garnet amphibolite and amphibolite. It was initially metamorphosed to lower amphibolite facies and was later retrograded to lower greenschist facies. The lower amphibolite facies metamorphism is Neoproterozoic (670 to 650 Ma) as inferred from the ⁴⁰Ar/³⁹Ar high temperature release spectra of hornblende separates. A U-Pb zircon age of 646.7 ± 2.7 Ma obtained for a pegmatite that intruded deformed rocks is taken to be the minimum age of the Formation.

The platformal Islesboro Formation probably unconformably overlies the Seven Hundred Acre Island Formation. It is primarily turbiditic pelite with many beds of quartzite, impure dolomitic marble, some conglomerate, and a few feldspar-rich volcanoclastic beds and is thought to be either Neoproterozoic or Cambrian. It was metamorphosed only to lower greenschist facies, possibly in the same event that retrograded the Seven Hundred Acre Island Formation.

Geochemical interpretations of minor and trace element analyses of six amphibolite and four schist samples from the Seven Hundred Acre Island Formation show that the protoliths of the amphibolite samples were intermediate between tholeiitic and within-plate type basaltic flows or dikes that intruded attenuated continental crust, or were eroded from these basalts. Four amphibolite and three schist samples analyzed for Pb isotopes were found to be enriched in radiogenic Pb. The Pb isotopic compositions are similar to those in peri-Gondwanan basement rocks from Atlantic Canada.

The peri-Gondwanan Islesboro block was placed against the peri-Gondwanan Middle and Late Cambrian Ellsworth terrane on the east by significant Late Silurian strike-slip faulting. The Late Silurian or Early Devonian Turtle Head dextral strike-slip fault separates the Islesboro block from the peri-Gondwanan St. Croix terrane to the west. The timing and nature of the movements of these faults are given from offsets of the isograds around the zircon-dated Late Silurian Sedgwick Granite (419.5 ± 1.4 Ma) and South Penobscot Intrusive Suite (419.2 ± 2.2 Ma). These terranes and others like them in Atlantic Canada make up a composite terrane that is different from the composite Avalonian terranes *sensu stricto* of southeastern New England and Atlantic Canada. The composite peri-Gondwanan terrane that included the Islesboro block was amalgamated with the margin of ancestral North America in the Silurian.

Deux séquences stratigraphiques en plates-formes sont présentes à Islesboro, dans la baie Penobscot, au Maine. La Formation âgée de Seven Hundred Acre Island a au moins 200 mètres d'épaisseur; sa base n'affleure pas; et elle comporte des blocs limités par des failles de marbre dolomitique siliceux à rubans de couleur, du quartzite riche en muscovite, du schiste à muscovite-grenat-stauroilite-andalousite resplendissant à gros grain et des métapelites calcaires, avec une quantité restreinte d'amphibolite et d'amphibolite à grenat. Elle avait initialement subi un métamorphisme jusqu'au faciès amphibolitique inférieur et elle a plus tard régressé à un faciès de schistes verts bas. Le métamorphisme du faciès amphibolitique inférieur est survenu au cours du Protérozoïque supérieur (670 à 650 Ma) d'après ce qu'on peut déduire des spectres de libération à haute température ⁴⁰Ar/³⁹Ar des éléments de la hornblende. L'âge de 646,7 ± 2,7 Ma obtenu par datation au U-Pb à partir de zircon dans le cas d'une pegmatite faisant intrusion dans les roches déformées est considéré comme l'âge minimal de la Formation.

La Formation en plate-forme d'Islesboro recouvre probablement de façon discordante la Formation de Seven Hundred Acre Island. Cette formation est principalement constituée de pélite turbiditique comportant de nombreuses strates de quartzite, de marbre dolomitique impur, d'une certaine quantité de conglomérat et de quelques strates volcanoclastiques riches en feldspath. On pense qu'elle remonte au Protérozoïque supérieur ou au Cambrien. Elle s'est seulement métamorphisée jusqu'à un faciès de schistes verts bas, possiblement au cours des mêmes circonstances que celles où a régressé la Formation de Seven Hundred Acre Island.

Une interprétation géochimique des analyses des éléments mineurs et des éléments traces de six amphibolites et de quatre schistes de la Formation de Seven Hundred Acre Island révèle que les protolithes des amphibolites ont constitué des stades intermédiaires entre les écoulements ou les filons intrusifs basaltiques tholéiitiques et intra-plaques ayant fait

intrusion dans la croûte continentale atténuée, ou qu'ils proviennent de l'érosion de ces basaltes. On a constaté que quatre amphibolites et trois schistes dont on a analysé les isotopes de plomb étaient enrichis en plomb radiogénique. Les compositions isotopiques du plomb sont semblables aux roches du socle périgondwanien des provinces atlantiques.

Le bloc périgondwanien d'Islesboro a été placé contre le terrane périgondwanien du Cambrien moyen et supérieur d'Ellsworth à l'est par la formation de failles décrochantes prononcées au cours du Silurien supérieur. La faille décrochante dextrale du Silurien supérieur ou du Dévonien inférieur de Turtle Head sépare le bloc d'Islesboro du terrane périgondwanien de St. Croix à l'Ouest. Le moment et la nature des mouvements de la formation de ces failles nous sont fournis par les décalages des isogrades obtenus à proximité du granite du Silurien supérieur (situé par datation à partir de zircon à $419,5 \pm 1,4$ Ma) ainsi que la suite Penobscot intrusif du Sud ($419,2 \pm 2,2$ Ma). Ces terranes et d'autres comme eux dans les provinces de l'Atlantique constituent un terrane composite différent des terranes composites avaloniens, au sens strict, présents dans le sud-est de la Nouvelle-Angleterre et dans les provinces de l'Atlantique. Le terrane composite périgondwanien qui comprend le bloc d'Islesboro a été fusionné à la marge de l'Amérique du Nord ancestrale au cours du Silurien.

Traduit par la rédaction

INTRODUCTION

The Islesboro block, not called a terrane because of its small size, is mostly located in the town of Islesboro, a 25 km long and 3 km wide group of islands in Penobscot Bay, mid-coastal Maine (Fig. 1). The block extends northeastward under the bay to form an elongate block about 7 km long and 3 km wide on the mainland in the town of Castine. This fault-bounded block and its regional setting between the Ellsworth terrane on the east and the St. Croix terrane on the west is shown on Fig. 1, and its bedrock geology and the localities for the analyzed samples are shown on Fig. 2. Reports on various aspects of the geology and geophysics of the block were given by Stewart (1974), Stewart and Lux (1988), Stewart *et al.* (1993, 1995a, 1995b, 1998). Seismic reflection data (Stewart 1998, Fig. 4A) indicate that the Islesboro block and the adjacent terranes are themselves all allochthonous above the east-dipping west-directed Sennebec Pond thrust fault, which beneath Islesboro lies at depths of 7 to 8 km.

The purpose of this paper is to provide a description of the lithostratigraphy, metamorphism, geochemistry, and geochronology of rocks that were studied, and use these to interpret the tectonic setting of deposition, geologic history, and regional geological correlations of rocks in the Islesboro block. The geological time scales used are those of Tucker *et al.* (1998), Tucker and McKerrow (1995), and Plumb (1991).

ANALYTICAL METHODS

Petrographic analysis

Forty-five samples were used for petrographic examination. In addition, minerals in seven samples, including five that were chemically analyzed, were analyzed by Stewart and M.J.K. Flohr in 1985–86 with an SEMQ electron microprobe that was calibrated using analyzed minerals for internal standards. Iron analyzed was calculated as FeO.

Major, minor and trace element analyses

Ten samples from the Seven Hundred Acre Island Formation, and one from the muscovite pegmatite on Spruce Island that cuts it, were analyzed for major and trace elements.

Major and trace element data were obtained at the analytical laboratories of the U.S. Geological Survey using standard techniques summarized by Baedecker (1987).

Whole-rock Pb isotopic compositions

A subset of seven whole-rock samples, three of schist and four of amphibolite, was analyzed for Pb isotopic compositions. Whole-rock Pb isotopes were analyzed at the isotope laboratory of the U.S. Geological Survey, Reston, Virginia, using procedures previously outlined by Ayuso *et al.* (1996).

U-Pb isotopic analysis of zircon

Two- to eight-kilogram samples for isotopic dating were collected from natural exposures. Following pulverization in a jaw-crusher and rotary disc-mill, accessory minerals were extracted using established techniques of Wilfley table concentration and separation by heavy-liquid and magnetic susceptibility techniques. Zircon was selected from the high-purity mineral separates by handpicking with the aid of a binocular microscope. Faceted prismatic zircon or fragments with external faces were selected, air-abraded, and cleaned as is standard treatment for isotope dilution analysis.

Following a brief (10–15 min) wash in warm 2N HNO₃, a final selection of grains was made in ethanol. The selected grains were cleaned in warm, distilled 4N HNO₃ and acetone. Zircon fractions and single-crystals were spiked with a mixed ²⁰⁵Pb-²³⁵U tracer solution and dissolved in distilled 48% HF and concentrated HNO₃ in sealed TFE Teflon digestion vessels at 210°C for 72 hours. Isolation and purification of Pb and U followed the procedures described in Krogh (1973). Total procedure blanks during the course of these analyses were 1–3 pg for Pb and less than one pg for U.

Pb and U were loaded on single rhenium filaments with silica gel and phosphoric acid in a clean-air environment; isotopic measurements were made with a Fissons Sector-54 thermal ionization mass spectrometer equipped with eight Faraday detectors and an ion-counting Daly detector. Because of the small amount of Pb loaded for analysis (in some cases less than 100 pg), most analyses were performed using peak jumping on the Daly detector. Daly mass bias was periodically

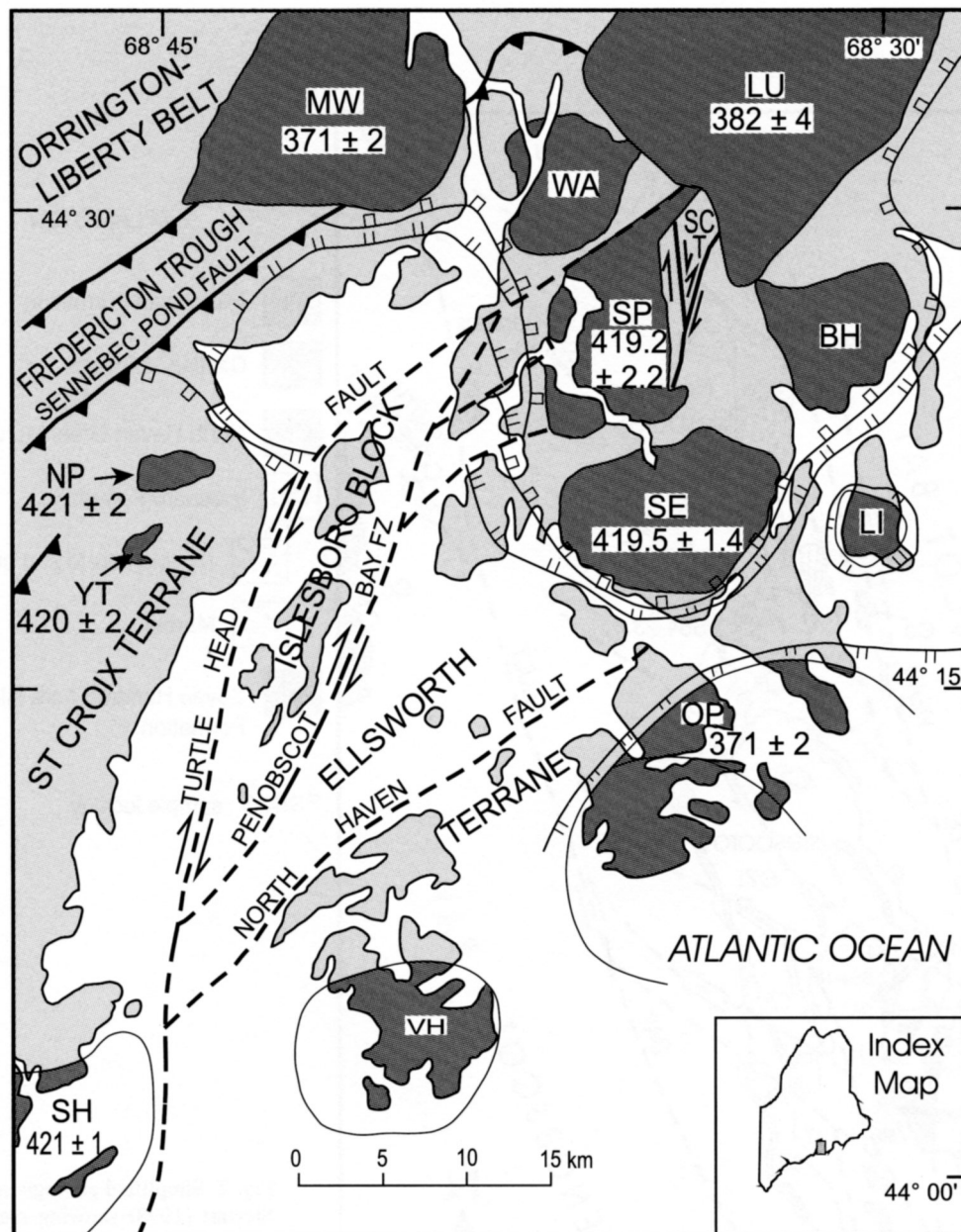


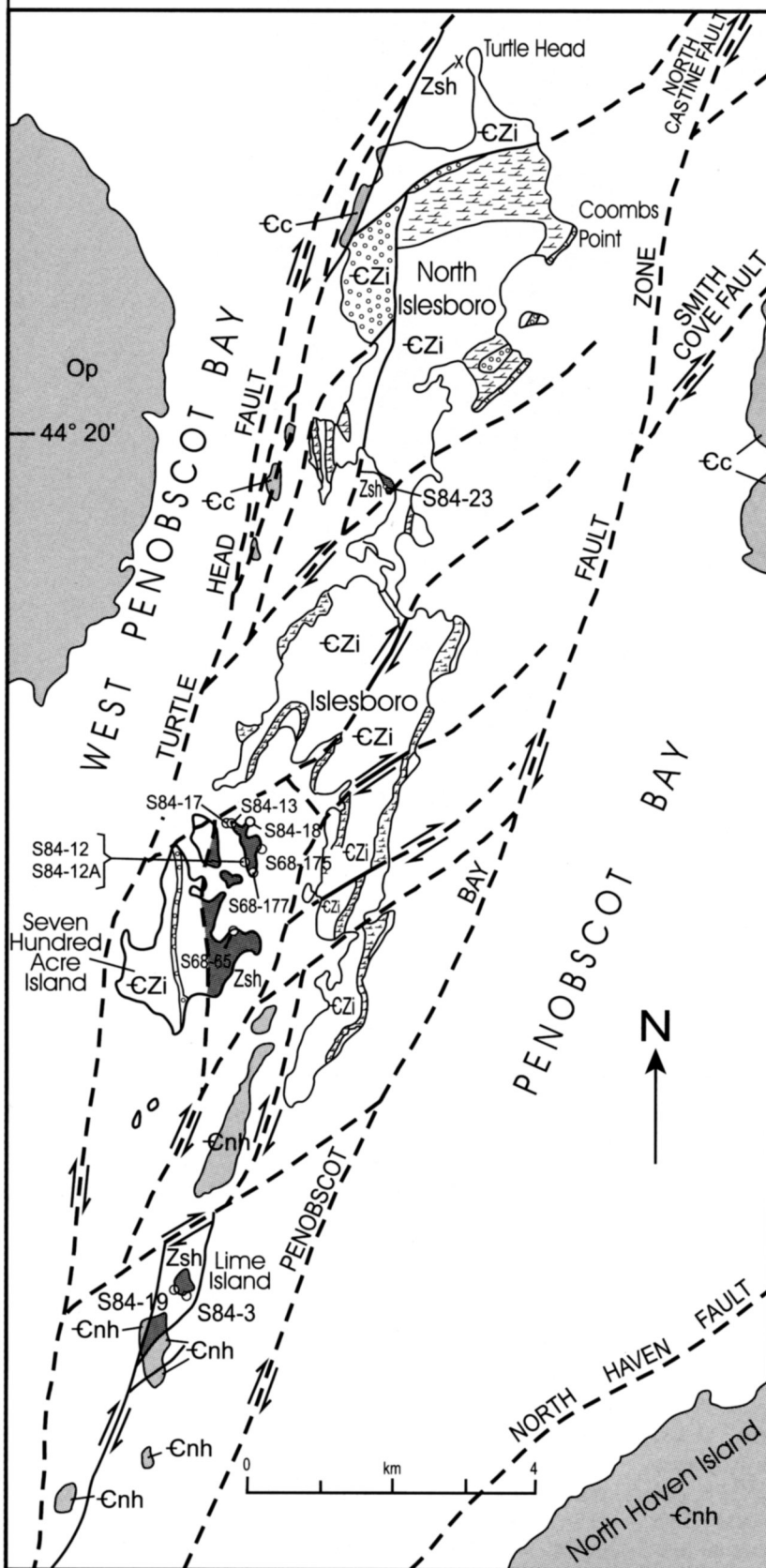
Fig. 1. Simplified geologic map of northern Penobscot Bay, Maine, showing the location of the Islesboro block and adjacent Ellsworth and St. Croix peri-Gondwanan terranes, the principal faults, metamorphic isograds, and plutons after Stewart (1998). The ages of the plutons are from Stewart (1998) and Tucker *et al.* (2001). Isograds: double barbed decoration = biotite, block decoration = andalusite. Abbreviations for plutons: BH, Blue Hill; LI, Long Island; LU, Lucerne; MW, Mount Waldo; NP, Northport; OP, Oak Point; SE, Sedgwick; SH, Spruce Head; SP, South Penobscot; VH, Vinalhaven; WA, Wallamatogus Mountain; YT, Youngtown. SCT = fault block of St. Croix terrane. More details of the Islesboro block are shown in Fig. 2.

monitored with National Institute of Standards and Technology and Central Bureau for Nuclear Measurements isotopic reference materials, and correction factors (and uncertainties) for Daly gain were used in data reduction. Mass-dependent fractionation corrections for Pb ($0.094 \pm 0.025\%$ amu⁻¹ at 1σ) and for U ($0.111 \pm 0.02\%$ amu⁻¹ at 1σ) were applied. For samples with total radiogenic Pb in excess of 1 ng, simultaneous measurements of ²⁰⁶Pb, ²⁰⁷Pb, ²⁰⁸Pb, and ²⁰⁵Pb were made in Faraday detectors, and ²⁰⁴Pb and ²⁰⁵Pb were measured in the Daly detector. In optimum cases an ion-beam between 0.5 and 1.5×10^{-13} A was maintained for ²⁰⁶Pb








(UO_2^+) in the Daly detector with an average ²³⁵U ion beam intensity of 2.5×10^{-13} A.

Measured ²⁰⁶Pb/²⁰⁴Pb ratios range between 658 and 16,364, reflecting the low concentration of Pb in laboratory reagents and materials and the low concentration of common Pb in zircon. The isotopic composition of the laboratory Pb is estimated from replicate measurements of reagents and total-procedure blanks determined since October 1994; values ($\pm 2\sigma$) are for the ²⁰⁶Pb/²⁰⁴Pb = $19.4 \pm 2.5\%$, ²⁰⁷Pb/²⁰⁶Pb = $0.804 \pm 2.2\%$, and ²⁰⁸Pb/²⁰⁶Pb = $2.0 \pm 2\%$. Initial-Pb compositions are calculated from the Pb-evolution model of Stacey and Kramers (1975) at the estimated age of the rock.

68° 00'



EXPLANATION

-  Penobscot Formation
-  Castine Volcanics
-  North Haven Greenstone
-  Islesboro Formation
-  Conglomerate and quartzite
-  Marble
-  Seven Hundred Acre Island Formation

○ S84-3 sample locality

Fig. 2. Simplified geologic map of Islesboro after Stewart (1998) showing mapped fault blocks of Seven Hundred Acre Island Formation, Islesboro Formation, North Haven Greenstone, and Castine Volcanics. Locations are shown for the analyzed samples from the Seven Hundred Acre Island Formation and muscovite granite. Modes of the analyzed samples are given in Appendix A.

The error of the initial-Pb compositions is assumed to be 4% (2σ). In all the cases reported here, even large variations in initial-Pb compositions result in insignificant changes to the calculated ages. Errors in $^{238}\text{U}/^{206}\text{Pb}$, $^{235}\text{U}/^{207}\text{Pb}$, and $^{207}\text{Pb}/^{206}\text{Pb}$ were calculated with the algorithm of Ludwig (1980) using a QuickBasic program written by L. Heaman and D.W. Davis of the Royal Ontario Museum (Toronto).

Because most of the zircon analyses cluster within analytical error of concordia, the cited age is the average of the $^{207}\text{Pb}/^{206}\text{Pb}$ ages, weighted according to the inverse variance of the individual ages (Ludwig 1992); the age error is simply the standard error of the average value calculated using the assigned error for each analysis. The reproducibility of the isotopic ages can be evaluated by the mean square of the weighted deviates (MSWD) which is a measure of the observed scatter to that predicted by the assigned errors to each data point. In two of the three samples, the MSWD is much less than one, indicating that the errors assigned to the individual $^{207}\text{Pb}/^{206}\text{Pb}$ ages may be somewhat over-estimated. All uncertainties are quoted, and shown in the concordia figures, at the 95% confidence limit.

$^{40}\text{Ar}/^{39}\text{Ar}$ release spectra of hornblende

The hornblende separates and flux monitors (SGB-7; muscovite, age 240.9 Ma relative to Mmhb-1) were wrapped in Sn foil, sealed in silica glass vials and irradiated in the L67 facility of the Ford Nuclear Reactor at the University of Michigan. The irradiation vessel was monitored such that all of the unknowns were within 10 mm of a flux monitor. Purification of the gases extracted from the flux monitors and unknowns followed standard procedures and the extraction-purification system is on line to the mass spectrometer. The isotopic composition of Ar was measured using a Nuclide 6-60-SGA 1.25 mass spectrometer at the University of Maine-Orono. Data reduction procedures are described in detail by Lux (1986). Ages were calculated using constants recommended by Steiger and Jager (1977). Analytical uncertainties associated with individual steps are at the 2σ level. Uncertainties associated with plateau ages are 2σ and include a 0.5% uncertainty in the J-value. The total gas ages represent weighted averages based on the amount of ^{39}Ar in each out-gassed step.

LITHOSTRATIGRAPHY

Seven Hundred Acre Island Formation

The Seven Hundred Acre Island Formation was mapped in three fault-bounded blocks, the largest 5 km by 2 km and the smallest less than 0.5 km by 0.5 km, juxtaposed by faults with the Islesboro Formation and Middle Cambrian North Haven Greenstone (Fig. 2). A fourth area about 5 m by 10 m was found on the west side of Turtle Head at low tide (Fig. 2). Thin- to thick-bedded, foliated, brown quartzite and muscovitic, sulphidic, and (or) feldspathic quartzite are interbedded with light-grey or buff, siliceous or graphitic, dolomitic marble, coarse-grained, splendent, muscovite-garnet-staurolite-andalusite schist, black hornblende-garnet

amphibolite, greenish- or brownish-grey, feldspathic, biotite-, phlogopite-, or idocrase- (pseudomorphs only) gneiss, and some thin, grey to black, graphitic, pelitic beds. Bedding is 1 cm to 0.5 m thick and forms aggregate packages of beds that are 2 to 15 m thick. The formation is at least 200 m thick but it is complexly folded and faulted so that its true thickness is not known. The base of the formation is not exposed and the basement upon which it was deposited is unknown. The type section is on eastern Seven Hundred Acre Island along the south shore of Cradle Cove and on the south shore of the adjacent Spruce Island in Gilkey Harbour (Stewart 1998). Dark, greenish-black amphibolite occurs throughout the section as dikes and sills of metabasalt, as weakly pillowed flows, or as metamorphosed sediments derived from basalt. Lime Island, in the southern block of exposed Seven Hundred Acre Island Formation, is made up of tightly folded and faulted dolomitic marble with dikes of metabasalt.

The protoliths of the metamorphosed sedimentary rocks of the Seven Hundred Acre Island Formation are interpreted to have been deposited as mature sediments in a platformal sequence. Some protoliths were basalts that were erupted as flows or were intruded as dikes into the Formation or were sediments derived from them.

The age of deposition of the formation is not known; the estimate by Brookins (1976) of this age from a whole-rock Rb-Sr isochron age is unreliable. The formation is metamorphosed to lower amphibolite facies and is deformed. The age of post-lower amphibolite facies cooling is estimated to be 670 to 650 Ma based on $^{40}\text{Ar}/^{39}\text{Ar}$ release spectra from hornblende, as described below in the section on geochronology. Muscovite pegmatites intrude the deformed and metamorphosed formation so their age is the same or younger than the age of metamorphism. A muscovite pegmatite on Spruce Island yielded zircon with an age of 647 ± 2.7 Ma, as described below.

Islesboro Formation

The sedimentary protoliths of the Islesboro Formation have been metamorphosed to lower greenschist facies. The Islesboro Formation is predominantly light grey to greenish-grey, fine-grained metasilstone interbedded with quartzite that may be feldspathic or calcareous. Some beds are light purplish-grey. Beds range in thickness from 2 cm to 20 cm and have distinct bases and gradational tops. Several stratigraphic successions of quartzite and fine-grained, siliceous, dolomitic marble are shown on Fig. 2 from Stewart (1998). Beds of marble weather rusty brown or dark brown but are light- to medium-grey or bluish grey when freshly broken. Individual beds vary in thickness from 2 cm to 2 m and an aggregate package of beds may be up to 6 m thick. A carbonate unit at Coombs Point was called the Coombs Limestone member by Smith *et al.* (1907). However, several sequences of sandy and micaceous thinly bedded marble and dolomite occur in the Islesboro Formation (Fig. 2) and cannot all be the same unit, so this member has been abandoned.

Conglomerate sequences up to 20 m thick are interbedded with quartz-rich metasandstone and metasilstone in different parts of the overall stratigraphic sequence. Beds of quartzite

and metaconglomerate with subrounded quartzite or quartz pebbles and cobbles with a sandy matrix typically are 0.5 to 1 m thick but may be up to 3 m thick. Some of these beds contain grey chert or red jasper pebbles. Quartz pebble metaconglomerate interbedded with greenish- and purplish-grey metapelite and quartzite forms a band extending from the south to the north across Seven Hundred Acre Island. A 10 m-thick conglomerate sequence overlying dolomitic marble near Coombs Point and Coombs Cove was called Battie Quartzite by Smith *et al.* (1907), but the Battie member is now known to be part of a quite different lithologic assemblage in the St Croix terrane west of Penobscot Bay (Berry and Osberg 2000).

Some beds rich in plagioclase and chlorite, and containing some lithic clasts possibly originally derived from volcanic wacke, occur on the east side of Turtle Head and in North Islesboro along Main Road. An albitized felsic volcanoclastic bed 2 m thick occurs in the borrow pit on Kedears Hill, North Islesboro.

The overall thickness of the formation is unknown because of faulting but probably exceeds 700 m. Mafic dikes, as well as a few dacite dikes also metamorphosed to lower greenschist facies, occur locally. Islesboro Island is divided into blocks by right-lateral strike-slip faults that are subparallel to large, regional, strike-slip faults both east and west of the island, as shown by Figs. 1 and 2. Most of these blocks are composed of Islesboro Formation except in the southeasternmost part of the town where fault blocks of North Haven Greenstone occur adjacent to both Islesboro and Seven Hundred Acre Island formations. Stratigraphy within the Islesboro Formation can be mapped for kilometres within a fault block, and from one fault block to another (Fig. 2). Folds in the Islesboro Formation have amplitudes and wavelengths of 0.5 km to at least 2.0 km, and many associated minor folds have axial surfaces with similar northeasterly and northern trends. Detailed studies of the folding and spectacular spaced cleavage developed in graded beds at Turtle Head, Islesboro, were published by Gregg (1985, 1986). Folding in the Islesboro Formation is not related to folding in the older Seven Hundred Acre Island Formation.

The Islesboro Formation is a cover sequence of platformal or shallow water slope-rise metasedimentary rocks. Little-deformed beds that are highly variable in thickness probably represent channel fillings. Conglomerate beds composed mostly of quartz pebbles and quartzite cobbles are both normally and reversely graded and were deposited in channels and outwash plains. The Islesboro Formation is interpreted to have been deposited unconformably upon basement of the Seven Hundred Acre Island Formation. However, the unconformity has not been observed, and the Islesboro Formation is separated by faults from basement. These faults are not metamorphic isograds. No evidence of metamorphic minerals or lithic fragments derived from the Seven Hundred Acre Island Formation was found in the metasedimentary rocks in the Islesboro Formation.

Late Silurian granitic plutons

The timing of movement on the faults that bound the

Islesboro block is constrained by the ages of the Sedgwick Granite and the South Penobscot Intrusive Suite whose contact metamorphic aureoles are cut by these faults. The plutons do not occur in the Islesboro block itself, but lie near its northeastern end in rocks of the Ellsworth terrane (Fig. 1; Stewart 1998). The ages of these plutons have been determined using U-Pb on zircons yielding Late Silurian ages, as described in the section on geochronology.

The Sedgwick Granite is a grey, medium- to coarse-grained, equigranular, biotite granite that contains 35% microcline, 32% quartz, 28% oligoclase, 4% biotite, and trace amounts of zircon, apatite, and titanite. The sample for U-Pb dating was collected along route 172, 200 m north of North Sedgwick. Isograds of the contact metamorphic aureole of this pluton are offset by the Penobscot Bay-Smith Cove fault.

The South Penobscot Intrusive Suite is a crudely normally zoned intrusion containing a variety of plutonic rocks that range from massive or layered gabbro to quartz diorite, tonalite, granodiorite, and granite (Stewart *et al.* 1988). The dated sample (S70-89) is porphyritic Grey Ridge Granite. It is light-grey, coarse-grained, biotite-muscovite granite containing approximately 20% microcline phenocrysts, and is from the youngest part of the Intrusive Suite (Stewart *et al.* 1988). The dated sample contains 37% quartz, 33% microcline, 26% oligoclase, 4% biotite and traces of zircon, apatite, titanite, and muscovite.

METAMORPHISM

Seven Hundred Acre Island Formation

The metamorphic mineralogy of the Seven Hundred Acre Island Formation is complex because of the presence of many varieties of metasedimentary rocks and some metabasalt. Rocks with lower amphibolite facies assemblages were faulted and subsequently retrograded to lower greenschist facies. The amount of retrograde metamorphism varies from relatively minor with preserved pseudomorphs of garnet or idocrase, to extensive involving the nearly complete recrystallisation of the lower amphibolite facies assemblage to a lower greenschist assemblage. Rocks retaining the effects of the lower amphibolite facies metamorphism are described first, according to rock type by groups of similar composition or metamorphic assemblage. Effects of the younger lower greenschist facies metamorphism are then described for the same rock types.

Lower amphibolite facies assemblages

Mapped deformed basaltic dikes and some basaltic flows with relict pillow structures have been metamorphosed to amphibolite. The lower amphibolite facies assemblage is hornblende-plagioclase-titanite-magnetite, but these rocks are most commonly extensively retrograded. Two (S84-12, S84-13) of three metabasalts chemically analyzed mostly retain the lower amphibolite assemblage; the third, S84-19, is mostly retrograded to actinolite and chlorite. Hornblende from samples S84-12 and S84-13 (Table 1) has slightly greater amounts of ferrohastingsite than pargasite components

Table 1. Representative electron microprobe analyses of metamorphic minerals from the Seven Hundred Acre Island Formation

Sample #	SiO ₂	TiO ₂	Al ₂ O ₃	FeO	MnO	MgO	CaO	Na ₂ O	K ₂ O	Total	TSi	TAI	CAI	CFe ³	CMg	CTi	BMn	BFe ²	BCa	BNa	ANa	AK	ACa	Fe#		
Amphibole analyses and structural formulae (based on 23 oxygen)											TSi	TAI	CAI	CFe ³	CMg	CTi	BMn	BFe ²	BCa	BNa	ANa	AK	ACa	Fe#		
Hornblende																										
S84-12	41.16	0.91	13.12	19.55	0.26	8.32	11.49	1.72	0.41	98.94	6.44	1.55	0.75	2.30	1.85	0.10	0.33	0.14	1.83		0.50	0.09	0.01	0.57		
S84-12A	42.82	0.99	14.17	19.32	0.19	7.84	11.59	1.80	0.43	99.15	6.37	1.63	0.85	2.30	1.77	0.11	0.02	0.11	1.85	0.23	0.52	0.18		0.58		
S84-13	45.09	1.07	9.03	18.82	0.06	10.79	11.09	2.14	1.18	99.27	6.72	1.28	0.31	2.17	2.40	0.12	0.01	0.18	1.77		0.62	0.23		0.49		
Actinolite																										
S84-12	53.07	0.02	1.39	16.73	0.13	13.25	12.75	0.15	0.05	97.54	7.81	0.19	0.05	2.04	2.91	0.00	0.02	0.02	1.97		0.04	0.01	0.04	0.41		
S84-13	52.67	0.07	1.49	14.86	0.06	14.97	12.36	0.45	0.15	97.08	7.73	0.26	0.00	1.71	3.28	0.01	0.01	0.12	1.81		0.13	0.03	0.07	0.23		
Mica analyses and structural formulae (based on 11 oxygen)											TSi	TAI	YAl	YMg	YTi	YFe ²					ANa	AK	ACa	Fe#		
Muscovite																										
S68-65	45.44	0.03	35.88	1.69	n.d.	0.25	0.05	0.58	9.53	93.43	3.07	0.93	1.92	0.03	0.00	0.10					0.08	0.82	0.01	0.79		
S68-65	45.65	0.03	35.87	1.72	n.d.	0.13	0.04	0.59	9.37	93.45	3.08	0.92	1.93	0.02	0.00	0.10					0.08	0.81		0.88		
S68-133	47.23	0.23	33.72	1.29	n.d.	1.15	n.d.	0.50	10.08	99.20	3.17	0.83	1.83	0.12	0.01	0.07					0.07	0.86		0.39		
S68-133	47.05	1.55	31.46	1.74	n.d.	1.53	0.01	0.31	9.92	93.55	3.19	0.81	1.70	0.15	0.08	0.10					0.04	0.87		0.39		
Phengite																										
S68-133	49.54	0.04	28.86	1.77	n.d.	3.84	0.05	0.18	10.02	94.30	3.32	0.68	1.60	0.38	0.00	0.10					0.02	0.86		0.20		
S68-65	47.93	0.03	34.55	2.24	n.d.	0.12	0.02	0.62	8.78	98.80	3.19	0.81	1.90	0.01	0.00	0.12					0.08	0.75		0.91		
Chlorite analyses and structural formulae (based on 16 oxygen)											TSi	TAI	YAl	YMg	YTi	YFe ²	YMn				ANa	AK	ACa	Fe#		
S84-12	25.70	0.05	20.44	30.24	0.34	12.17	0.08	0.02	0.03	89.07	5.47	2.53	2.61	3.86	0.01	5.39	0.06				0.01	0.01	0.02	0.58		
S84-12A	25.60	0.03	20.25	31.59	0.39	10.83	0.03	0.01	0.06	88.79	5.51	2.49	2.65	3.48	0.01	5.69	0.07				0.01	0.01	0.02	0.62		
S84-13	26.72	0.02	19.36	27.26	0.08	15.73	0.05	0.06	0.06	89.32	5.58	2.42	0.34	4.90	0.00	4.76	0.01				0.03	0.02	0.01	0.49		
S68-65	26.14	0.08	22.89	33.24	0.19	6.65	0.04	n.d.	0.02	89.25	5.60	2.40	3.37	2.12	0.01	5.95	0.03				n.d.	0.01	0.01	0.74		
Garnet analyses and structural formulae (based on 12 oxygen)											TSi	TAI	AAI	AFe ³	ATi	BMg	BMn	BFe ²	BCa	Alm	Pyr	Gro	Spe	And		
S84-12A	37.59	0.07	22.20	28.10	2.66	2.49	9.47	0.01	n.d.	102.99	2.91	0.09	1.93	0.15	0.01	0.29	0.17	1.67	0.79	57	10	20	6	7		
S68-65	37.89	0.07	21.99	35.96	0.68	3.06	2.83	0.01	n.d.	102.49	2.96	0.04	2.00	0.03	0.01	0.36	0.04	2.33	0.24	78	12	7	2	1		
Chloritoid analyses and structural formulae (based on 12 oxygen)											TSi	TAI	YAl	YMg	YTi	YFe ²	YMn							Fe#		
S68-65	24.17	0.06	40.81	26.47	0.57	1.11	0.04	n.d.	n.d.	93.27	2.00	3.00	0.98	0.14	0.00	1.83	0.04							0.93		

Recalculations were using the program MINPET. T, C, B, and A, or Y, refer to the cation sites in the standard mineral formulas. Fe² and Fe³ refer to ferrous and ferric iron, respectively. Fe# = Fe/Fe+Mg; n.d. = not determined. Sample locations are shown on Fig. 2.

(Fe/Fe+Mg = 0.49 to 0.57), and has about 1% TiO₂. Plagioclase is present as either relict phenocrysts or interstitial grains, but most of the calcic plagioclase is replaced by zoisite, or as albite plus other greenschist minerals. Quartz content is less than 10%. Titanite and magnetite are common accessory minerals.

Amphibolite and garnet amphibolite derived from sedimentary protoliths are variable in composition. Samples mostly derived from basaltic debris have abundant hornblende and titanite, but also have much quartz and muscovite. Hornblende from sample S84-12A has a Fe/Fe+Mg ratio of 0.58, and contains almost 1% TiO₂, (Table 1). Plagioclase is retrograded to albite. This sample has 10% garnet crystals up to 6 mm in size, which are almandine-rich (Alm₅₅). Trace amounts of apatite and magnetite are also present.

Muscovite - garnet - andalusite - staurolite - biotite - plagioclase - quartz splendent schists are important constituents of the Seven Hundred Acre Island Formation. Three of these rocks (S68-65, S68-177, and S84-23) were chemically analyzed (Table 2) and the minerals for two (S68-65, S68-133) were analyzed by electron microprobe. These schists are conspicuous because rounded garnets 2 to 5 mm across constitute up to 30% of the rock. Layers of small quartz grains within the garnet might be relicts of beds in the protolith. The garnets are almandine-rich (Alm₇₈; S68-65, Table 1). Large muscovite crystals are prominent and are commonly highly sheared so that muscovite is bent. The muscovite has more than 0.85 Al in tetrahedral positions per 11 oxygen and contains about 2% FeO and 0.15% MgO in solid solution. Andalusite is seldom apparent in hand specimen and staurolite is rare even in thin section. Andalusite is nearly ideal composition but no staurolite was found for analysis in the polished sections.

Siliceous marble contains more than 20% calcite and a variety of metamorphic minerals including phlogopite, tremolite/actinolite, and graphite. One sample (S84-17) has 2 cm long square prisms of fine-grained sericite that are interpreted to be pseudomorphs of idocrase. Biotite-rich gneiss with quartz and up to 25% albitic plagioclase make up some of the metapelitic interbeds within the siliceous marble.

Lower greenschist facies assemblages

Amphibolite samples vary from rocks containing well-preserved lower amphibolite facies assemblages to nearly completely retrograded lower greenschist assemblages. In amphibolite samples from lower greenschist assemblages the most common mineral is actinolite with a lower Fe/Fe+Mg ratio than the hornblende it was derived from, low tetrahedral Al, and low alkali content (Table 1). Chlorite is common with a Fe/Fe+Mg ratio nearly the same as in the original hornblende. Titanite is more abundant in retrograde lower greenschist rocks than in lower amphibolite facies rocks. Albite (An_{<5}) is common and is associated with retrograde epidote.

Amphibolite inferred to be of meta-sedimentary origin has been strongly retrograded and has gneissic texture. Actinolite has replaced most of the hornblende in many rocks, and it has a lower Fe/Fe+Mg ratio than the hornblende. Chlorite with approximately equal amounts of ferroan and

magnesian components has replaced most of the garnet. Epidote and titanite are abundant, and in some rocks calcite is present in small amounts.

Almost all garnet in the former muscovite-garnet-andalusite-staurolite-schist has been retrograded to ferroan chlorite, with Fe/Fe+Mg = 0.72–0.74 (Table 1). Much of the andalusite and almost all staurolite in these rocks has been retrograded to pleochroic blue-green needles of chloritoid with Fe/Fe+Mg = 0.93. Some muscovite in fine-grained aggregates with chloritoid is phengitic with low aluminum and higher amounts of Fe or Mg in solid solution (Table 1).

Retrograde metamorphic effects in siliceous marble include the formation of sericite pseudomorphs after idocrase (S84-17), chlorite after phlogopite, and the formation of almost pure calcite as the carbonate mineral. Most biotite-rich felsic gneisses are partly to wholly retrograded to chlorite- and albite-rich rocks.

Islesboro Formation

Regional metamorphism of the Islesboro Formation attained only lower greenschist facies and the mineral assemblages contain abundant quartz, albite, chlorite, muscovite, tremolite in calcite-rich rocks, with minor pyrite and iron oxides, and occasional titanite and tourmaline grains. Chemical equilibration was attained in metamorphism, as no relict calcic plagioclase or biotite was observed in petrographic studies by D.B. Stewart and M.J.K. Flohr (unpublished data). The formation is biotite grade only in the very northeastern most part north of Castine (Fig. 1) where some greenish-brown biotite is present in the contact metamorphic aureole of the South Penobscot Intrusive Suite. The prograde lower greenschist facies metamorphism of the Islesboro Formation is the same grade as that superimposed on the Seven Hundred Acre Island Formation and is presumably of the same age. However, that age has not been determined.

GEOCHEMISTRY OF THE SEVEN HUNDRED ACRE ISLAND FORMATION

Major, minor and trace element analyses

The major, minor, and trace element compositions of amphibolite, hornblende gneiss, schist, and granite pegmatite are given in Table 2. The abundances of the alkali elements are presumed to be unreliable for classification and for comparison as these rocks all have undergone some degree of alteration. However, the tectonic setting of origin of the protoliths of the amphibolites can be determined on the basis of their Ti, Zr, Nb, and Y contents (Pearce and Norry 1979; Pearce 1996).

Spider diagrams normalized to N-MORB (Pearce 1996) illustrate compositional data for the amphibolite samples from the Seven Hundred Acre Island Formation. Normalized plots for most samples lack spiked spidergrams (Fig. 3a) or pronounced troughs for Nb and Ta like those typical of calc-alkaline rocks (Pearce and Norry 1979; Pearce 1996). Instead, the Nb and Ta variations illustrate very small troughs or are

Table 2. Major and trace element contents and Pb isotopic compositions of whole rocks from the Seven Hundred Acre Island Formation and pegmatite, southern Maine. See Appendix A for rock names and modes.

Sample	S84-3 Amph.	S84-12 Amph.	S84-12A Gt. Amph.	S84-13 Amph.	S84-19 Amph. Dk.	S68-175 Hb. Gn.	S68-65 Sch.	S68-177 Sch.	S84-23 Spl. Sch.	S84-17 Calc. Pel.	S84-18 Peg.
Wt %											
SiO ₂	47.70	43.70	49.20	48.30	48.30	46.50	61.40	60.60	61.50	25.50	84.30
Al ₂ O ₃	13.30	12.30	12.70	13.10	18.10	13.40	18.50	21.90	21.50	6.85	9.37
Fe ₂ O ₃ ^a	15.60	19.10	17.60	15.50	10.70	12.10	10.80	6.03	1.78	2.29	0.37
MgO	6.70	7.78	5.59	5.87	6.06	11.80	1.54	0.51	0.88	15.30	0.31
CaO	6.97	9.01	8.93	7.32	7.66	9.56	0.56	0.02	0.02	18.20	0.21
Na ₂ O	3.42	1.43	1.73	4.67	3.11	1.18	0.41	0.40	0.72	1.16	1.90
K ₂ O	0.90	0.34	0.36	0.68	1.63	1.63	2.57	5.77	6.27	1.71	1.92
TiO ₂	2.68	3.65	2.16	2.50	1.24	0.75	0.93	0.84	0.75	1.05	0.11
P ₂ O ₅	0.28	0.29	0.17	0.22	0.17	0.06	0.13	0.05	0.05	0.09	0.10
MnO	0.17	0.28	0.24	0.02	0.12	0.18	0.10	0.02	0.02	0.04	<0.02
CO ₂	-	-	-	-	0.24	0.01	0.01	0.01	0.01	-	-
S	0.03	-	-	-	0.03	0.01	0.32	0.01	0.03	-	-
Cl	0.08	-	-	-	0.02	0.05	0.01	0.01	0.03	-	-
F	0.08	-	-	-	0.07	0.05	0.05	0.12	0.09	-	-
H ₂ O ⁺	2.4	-	-	-	3.0	2.7	3.0	5.2	5.3	-	-
H ₂ O ⁻	0.63	-	-	-	0.60	0.44	0.70	0.44	0.56	-	-
LOI ^b	-	1.99	1.75	1.48	-	-	-	-	-	27.90	1.10
Subtotal	100.94	-	-	-	101.05	100.42	101.03	101.93	99.51	-	-
Less O=F	0.07	-	-	-	0.05	0.04	0.19	0.06	0.06	-	-
Total	100.9	99.9	100.4	99.7	101.0	100.4	100.8	101.9	99.5	100.1	99.7
Trace elements, ppm											
Rb	12	6	<2	17	43	77	100	214	232	25	54
Sr	125	115	187	109	320	99	85	53	48	83	36
Ba	217	40	53	45	354	333	454	565	738	86	329
Pb	5	5	-	-	5	5	5	5	5	-	-
La	11.70	15.40	10.10	24.80	12.50	5.46	38.50	31.10	29.20	4.13	5.77
Ce	28.10	36.30	22.90	53.05	25.40	13.00	76.10	57.80	59.00	9.10	10.60
Nd	17.7	26.3	12.0	26.3	14.8	8.7	31.4	24.1	22.4	4.0	5.5
Sm	5.18	7.60	4.76	6.38	3.84	2.20	6.87	4.64	4.61	1.18	1.24
Eu	1.72	2.40	1.48	1.95	1.21	0.75	1.31	0.96	0.91	0.29	0.34
Tb	1.10	1.58	0.93	0.98	0.62	0.42	1.00	0.62	0.55	0.22	0.23
Yb	3.61	6.02	3.44	3.19	2.18	1.45	3.08	3.13	2.26	0.87	0.92
Lu	0.50	0.76	0.50	0.47	0.33	0.22	0.43	0.45	0.35	0.13	0.11
Y	34	57	31	33	2	11	36	32	2	4	11
Zr	168	133	111	156	95	52	233	118	92	125	30
Hf	4.39	3.97	3.21	4.11	2.27	1.23	6.11	3.21	2.49	3.01	0.46
Nb	23	14	<10	21	18	<10	11	19	14	20	13
Ta	1.72	0.84	0.65	1.43	0.94	0.44	0.90	1.09	0.94	1.22	0.29
Th	1.71	1.54	1.05	1.90	1.27	0.72	11.50	12.32	2.65	4.50	1.71
U	0.36	0.55	0.24	0.62	0.25	0.20	2.76	1.28	1.1	2.23	0.64
Sc	43.5	78.1	45.0	42.4	33.2	32.1	17.7	23.8	20.8	10.5	16.0
Cr	132.4	44.3	47.9	29.4	100.4	963.0	124.0	136.8	109.0	67.4	9.5
Co	41.20	54.00	54.40	26.15	33.40	39.40	41.00	14.88	0.66	0.60	0.39
Ni	68	60	46	43	53	240	100	69	5	<24	<26
²⁰⁶ Pb/ ²⁰⁴ Pb	20.871	21.057	-	-	19.718	19.021	17.955	21.083	24.223	-	-
²⁰⁷ Pb/ ²⁰⁴ Pb	15.891	15.989	-	-	15.786	15.684	15.697	15.923	16.173	-	-
²⁰⁸ Pb/ ²⁰⁴ Pb	38.663	41.316	-	-	37.647	39.541	37.659	42.954	43.493	-	-

^a Total iron as Fe₂O₃, ^b Loss on ignition at 900°C

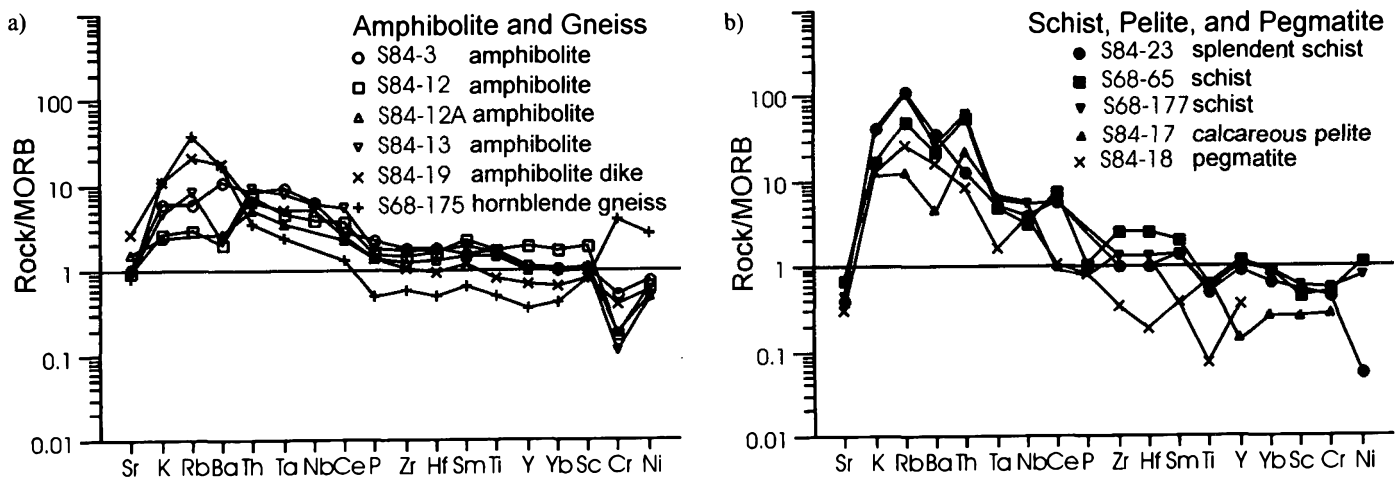


Fig. 3. Spider diagrams of a) amphibolite and hornblende gneiss samples and b) schist, metapelite, and pegmatite samples from the Seven Hundred Acre Island Formation normalized to N-MORB (Pearce 1996).

slightly enriched relative to the abundances of Th and Ce. The amphibolites lack troughs for Ti and Sc though they have moderate troughs for Cr, indicating fractionation of Fe-Ti oxides.

The normalized diagram for the schist samples (Fig. 3b) shows small troughs at Ba, Ta, Nb, Sr, and P (and Ti?) and a general enrichment for K, Rb, and Th. The small relative depletions in Nb and Ta relative to Th and Ce in these samples are compatible with an origin in an orogenic arc (Pearce and Norry 1979; Pearce 1996).

Chondrite-normalized variations for the REE show that the amphibolite samples (Fig. 4a) have flat REE patterns and no evidence for depletion of Eu by plagioclase fractionation or the heavy REE (HREE) by garnet fractionation. In contrast, the schist samples as a group are light REE (LREE) enriched (from about 10 \times to about 100 \times chondrite) and generally display smooth patterns (Fig. 4b). The HREE patterns are flat and some of the samples have small to moderate negative Eu anomalies.

Although the number of amphibolite samples in the study is limited, immobile trace element abundances (Fig. 3a) and their ratios (Th, Hf, Ta, Yb, etc.) as in Fig. 5 (after Wood

1980) are meaningful even if the rocks have been affected by metamorphic processes. Such trace element features can be used to estimate the original magmatic compositions of the mafic rocks and infer the nature of their sources and tectonic settings.

The schist samples have widely variable compositions and are insufficient in number for detailed geochemical interpretation. Their trace element abundances resemble those of granitic rocks from arc environments (e.g., Ta and Yb; Pearce *et al.* 1984).

Pb isotopes

Pb isotopic compositions of whole rocks from the Seven Hundred Acre Island Formation exhibit a wide range in $^{206}\text{Pb}/^{204}\text{Pb}$ from about 17.96 to 24.22, in $^{207}\text{Pb}/^{204}\text{Pb}$ from 15.68 to 16.17, and $^{208}\text{Pb}/^{204}\text{Pb}$ from 37.65 to 43.49 (Table 2). Age-corrected compositions (at 647 Ma) using measured U, Th, and Pb abundances are shown in Fig. 6. The Pb isotopic compositions are more radiogenic than the average Pb crustal evolution curve (Stacey and Kramers 1975) and attest to the contributions of evolved isotopic reservoirs in continental crust.

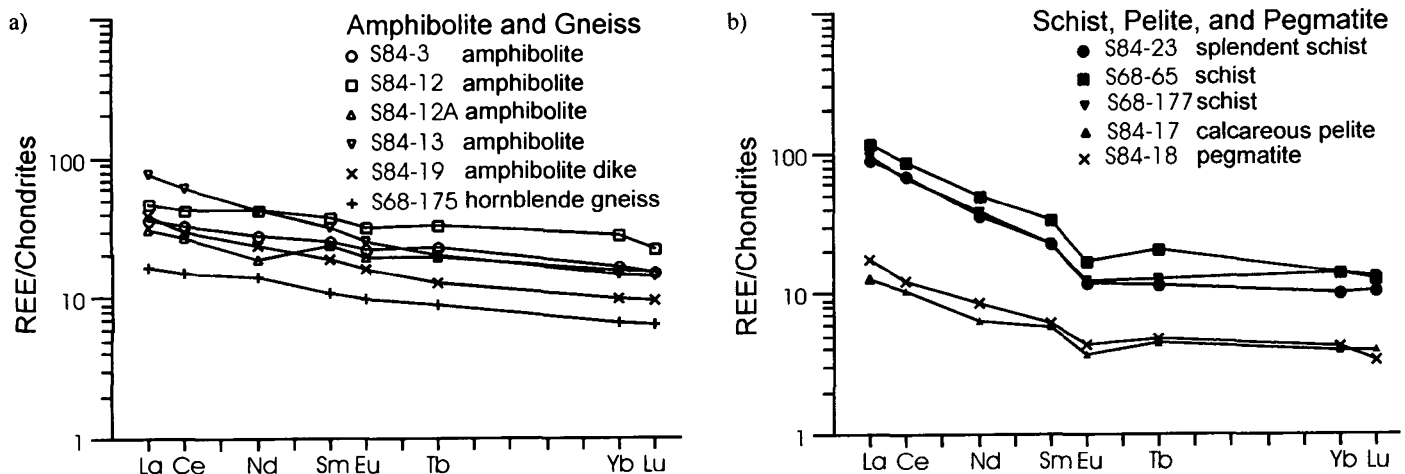


Fig. 4. Chondrite-normalized REE plots for a) amphibolite and hornblende gneiss samples and b) schist and metapelite samples from the Seven Hundred Acre Island Formation and from pegmatite that intruded it.

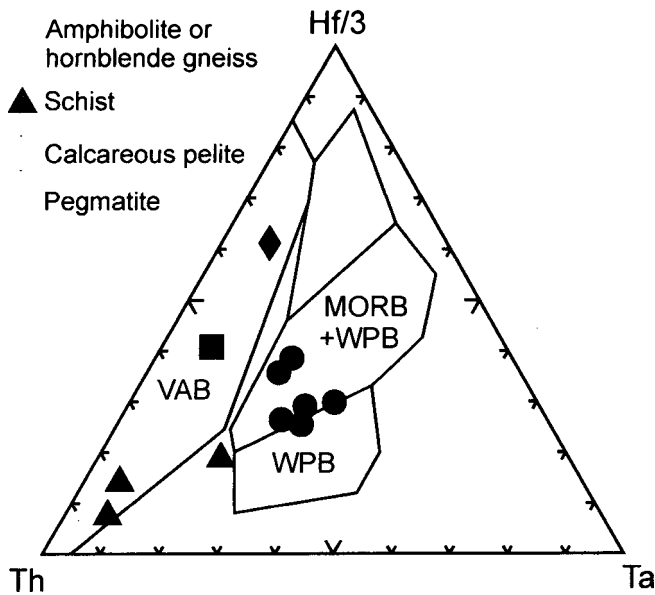


Fig. 5. Th-Hf-Ta diagram showing compositions of rocks from the Seven Hundred Acre Island Formation and from pegmatite that intruded it. The amphibolite samples have tectonomagmatic affinities resembling E-type MORB and within-plate tholeiite sources; they also overlap the field for alkaline within-plate basalts. Labelled fields from Wood (1980) and Pearce (1996): VAB, volcanic arc basalt; MORB, mid-ocean ridge basalt and within plate basalt (E-type MORB, enriched-type of mid-ocean ridge basalt); WPB, within-plate basalt.

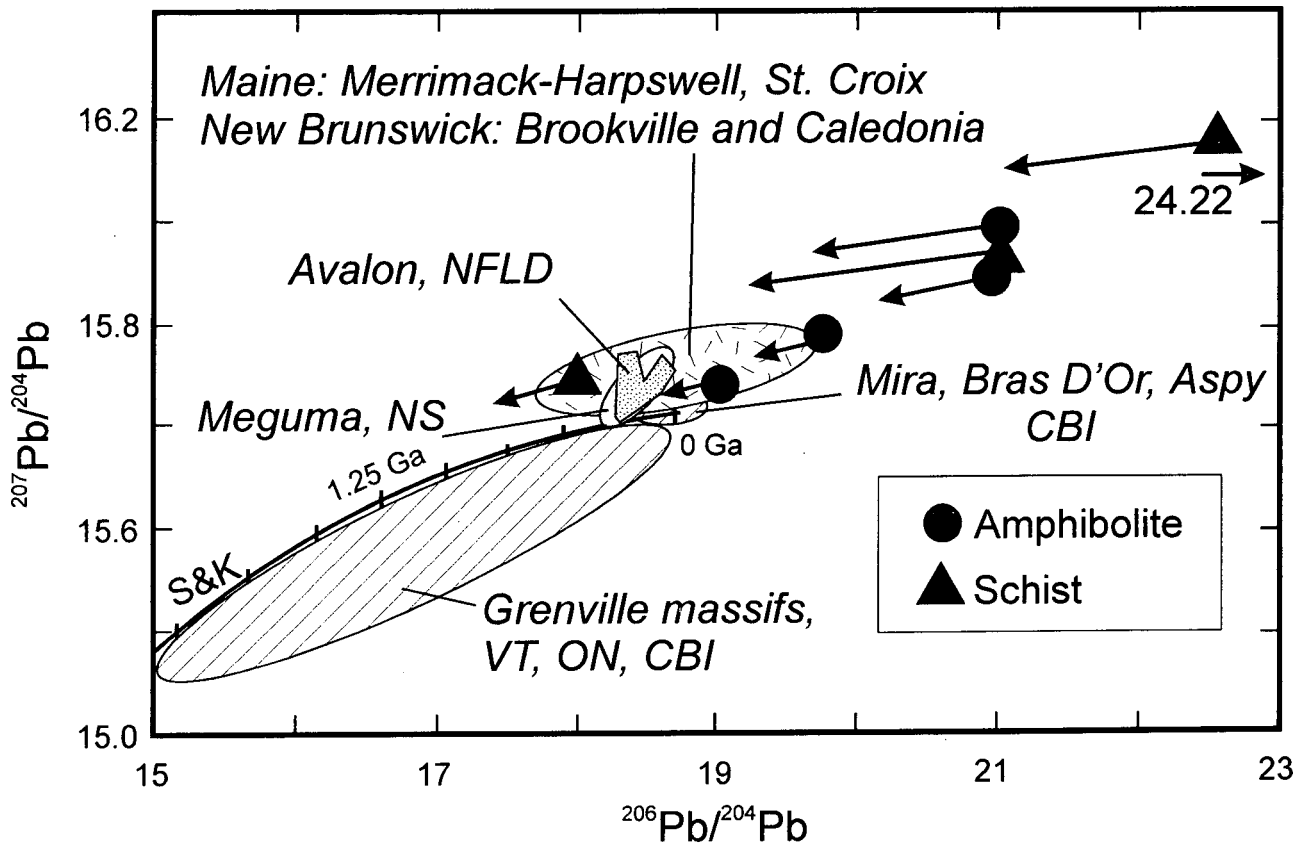


Fig. 6. $^{207}\text{Pb}/^{204}\text{Pb}$ vs. $^{206}\text{Pb}/^{204}\text{Pb}$ isotopic compositions of rocks from the Seven Hundred Acre Island Formation and comparisons to Grenville massifs and peri-Gondwanan basements from New England and Atlantic Canada. Vectors illustrate the direction and estimated change in Pb isotopic compositions of the amphibolite and schist from the Seven Hundred Acre Island Formation using measured U, Th, and Pb contents at 647 Ma. Sources of data: VT, Grenville basement rocks (gneisses and schists) from Vermont (R. Ayuso and N. Ratcliffe, unpublished data for Green Mountains Massif, 1996); ON, Ontario (DeWolf and Mezger 1994); CBI, Blair River complex (gneiss) in Cape Breton Island (Ayuso *et al.* 1996); peri-Gondwanan basement rocks in New Brunswick (Ayuso and Bevier 1991; Whalen *et al.* 1996a); granitic plutons and gneisses from the Mira, Aspy, and Bras d'Or blocks in Cape Breton Island (CBI) (Ayuso *et al.* 1996); peri-Gondwanan gneisses from the Brookville and Caledonia blocks from New Brunswick, the Merrimack-Harpswell and St. Croix blocks from Maine, and Paleozoic granitic rocks from the Northern Appalachians of New England and Atlantic Canada (e.g., Ayuso 1999; Ayuso and Bevier 1991; Ayuso and Arth 1997; Whalen *et al.* 1994a, b; 1996a, b); Devonian granitic rocks from the Canso area, Meguma block, Nova Scotia (NS) (R. Ayuso and J. Hill, unpublished data, 1994); granitic rocks from the Ackley batholith in the Avalon of Newfoundland (NFLD) (R. Ayuso, D. Kontak and D. Strong, unpublished data, 1994); crustal evolution curve (S&K, Stacey and Kramers 1975) showing tick marks at 0.25 Ga.

GEOCHRONOLOGY

U-Pb ages of zircon

Muscovite pegmatite at Spruce Island

The muscovite pegmatite cuts dolomitic limestone in metamorphic rocks at lower amphibolite grade at the northeast end of Spruce Island. The pegmatite is weakly cleaved but quite fresh. A whole rock chemical analysis of pegmatite sample S84-18 is given in Table 2 and a modal estimate is given in the appendix. Two zircon fractions for sample RT-537HA-5, one of 3 grains and one single grain, were analyzed (Table 3). One is 0.8% discordant with a $^{207}\text{Pb}/^{206}\text{Pb}$ age of 646 ± 3 Ma and the other is 2.9% discordant with a $^{207}\text{Pb}/^{206}\text{Pb}$ age of 649 ± 6 Ma. Uranium contents approach 2500 ppm so some Pb-loss arising from radiation damage is possible. If discordance arises from time integrated Pb-loss the $^{207}\text{Pb}/^{206}\text{Pb}$ age is the best estimate of the age of pegmatite emplacement. The weighted mean $^{207}\text{Pb}/^{206}\text{Pb}$ age of the two analyses is 646.7 ± 2.7 Ma (Fig. 7a). Estimates of the age of this pegmatite given in Brookins (1976) are incorrect.

Sedgwick Granite

A sample of granite of the Sedgwick Granite was collected along Route 172, 200 m north of North Sedgwick. Zircon from this sample, RT-93SP-17, is dominated by a population of clear long-prismatic faceted prisms. Three

fractions of long, prismatic grains were analyzed (Table 3). All are concordant within error and only moderate age discordance is observed (0.22–0.37%, Fig. 7b). Whereas the weighted-average of the $^{206}\text{Pb}/^{238}\text{U}$ ages is 418.3 ± 2.0 Ma, the weighted-average $^{207}\text{Pb}/^{206}\text{Pb}$ age is 419.5 ± 1.4 Ma (Fig. 7b). The difference between the two can be attributed to less than 0.4% of Pb-loss, and the $^{207}\text{Pb}/^{206}\text{Pb}$ age may therefore be the best estimate of the time of granite emplacement. Using the time-scale of Tucker and McKerrow (1995) a latest Silurian (Ludlow or Pridoli) emplacement age is implied. The weighted average $^{207}\text{Pb}/^{206}\text{Pb}$ is well defined (MSWD = 0.766) indicating that the assigned errors to the individual $^{207}\text{Pb}/^{206}\text{Pb}$ ages encompass the true age.

Grey Ridge Granite, South Penobscot Intrusive Suite

A sample (S70-89) of the Grey Ridge Granite, the youngest phase of the South Penobscot Intrusive Suite, yielded a zircon concentrate that had been previously studied by T.W. Stern (in Stewart *et al.* 1988, p. 90) to obtain an age of 424.4 ± 6 Ma. Two samples of clear and one ample of pale brown flat or long prisms and one sample of pale brown long prisms (Table 3) were selected from this concentrate for further study. The weighted mean $^{207}\text{Pb}/^{206}\text{Pb}$ age for these nearly concordant samples is 419.2 ± 2.2 Ma (Fig. 7c), and this is taken to be the age of the youngest part of the South Penobscot Intrusive Suite.

Table 3. U-Pb isotope dilution analyses.

No.	Fractions Properties (1)	Concentrations				Atomic Ratios					Ages $^{207}\text{Pb}/^{206}\text{Pb}$ (6)
		Wt. μg (2)	Pb* ppm (2)	U ppm (2)	Pb pg (3)	Th U (4)	$^{206}\text{Pb}/^{204}\text{Pb}$ (5)	$^{207}\text{Pb}/^{206}\text{Pb}$ (6)	$^{207}\text{Pb}/^{235}\text{U}$ (6)	$^{206}\text{Pb}/^{238}\text{U}$ (6)	
Spruce Island Pegmatite [RT-937HA-5]											
1	3 gr, pb,t-p	2	218	2286	5.2	0.002	5,863	0.06119 ± 0.00004	0.8819 ± 0.0014	0.10453 ± 0.00015	646.0 ± 1.5
2	1 gr, pb,t-p	3	246	2611	75.9	0.010	657.8	0.06130 ± 0.00009	0.8703 ± 0.0018	0.10296 ± 0.00016	649.9 ± 3.3
Sedgwick Pluton [RT-93S-17]											
3	3 gr, cl,c,l-p	10	11.6	171	1.1	0.376	6,437	0.05519 ± 0.00007	0.5109 ± 0.0008	0.06715 ± 0.00007	419.8 ± 2.8
4	20 gr, cl,c,l-p	69	15.6	230	4.2	0.375	16,050	0.05520 ± 0.00002	0.5106 ± 0.0007	0.06710 ± 0.00009	420.2 ± 0.9
5	11 gr, cl,c,l-p	33	13.3	198	1.6	0.361	16,364	0.05515 ± 0.00003	0.5085 ± 0.0006	0.06687 ± 0.00007	418.2 ± 1.3
South Penobscot Pluton [S-70-89]											
6	1 gr, pb,cr,l-p	3	139	2162	6.4	0.208	4,775	0.05522 ± 0.00006	0.5084 ± 0.0009	0.06678 ± 0.00011	421.1 ± 2.6
7	2 gr, pb,fl	7	144	2279	14.2	0.147	4,582	0.05517 ± 0.00004	0.5075 ± 0.0008	0.06672 ± 0.00009	419.2 ± 1.5
8	1 gr, cl,c,l-p	2	112	1683	3.0	0.332	5,026	0.05514 ± 0.00011	0.5068 ± 0.0012	0.06666 ± 0.00018	418.0 ± 4.3
9	1 gr, cl,c,fl	3	92.7	1453	3.0	0.163	5,363	0.05512 ± 0.00007	0.5094 ± 0.0009	0.06702 ± 0.00009	417.2 ± 2.8

Notes:

- (1) All analyses are of zircon. Cardinal number indicates the number of zircon grains analyzed (e.g. 7 grains); all zircon grains were selected from non-paramagnetic separates at 0° tilt at full magnetic field in Frantz magnetic separator; c = colorless; cr = cracked; cl = clear; fl = flat; l-p = long-prismatic; pb = pale brown; t-p = tips from prisms. All grains air-abraded following Krogh (1982).
- (2) Pb* = radiogenic Pb. Concentrations are known to $\pm 30\%$ for sample weights of about 30 μg and $\pm 50\%$ for samples $< 3 \mu\text{g}$.
- (3) Corrected for 0.0125 mole fraction common-Pb in the $^{205}\text{Pb} - ^{235}\text{U}$ spike.
- (4) Calculated Th/U ratio assuming that all ^{208}Pb in excess of blank, common-Pb, and spike is radiogenic ($\lambda^{232}\text{Th} = 4.9475 \times 10^{-11} \text{ y}^{-1}$)
- (5) Measured, uncorrected ratio.
- (6) Ratio corrected for fractionation, spike, blank, and initial common-Pb (at the determined age from Stacey and Kramers 1975). Pb fractionation correction = $0.094\% / \text{amu}$ ($\pm 0.025\%$, 1σ); U fractionation correction = $0.111\% / \text{amu}$ ($\pm 0.02\%$, 1σ). U blank = 0.2 pg; Pb blank ≤ 10 pg. Absolute uncertainties (1σ) in the Pb/U and $^{207}\text{Pb}/^{206}\text{Pb}$ ratios calculated following Ludwig (1980). U and Pb half-lives and isotopic abundance ratios from Jaffey *et al.* (1971).

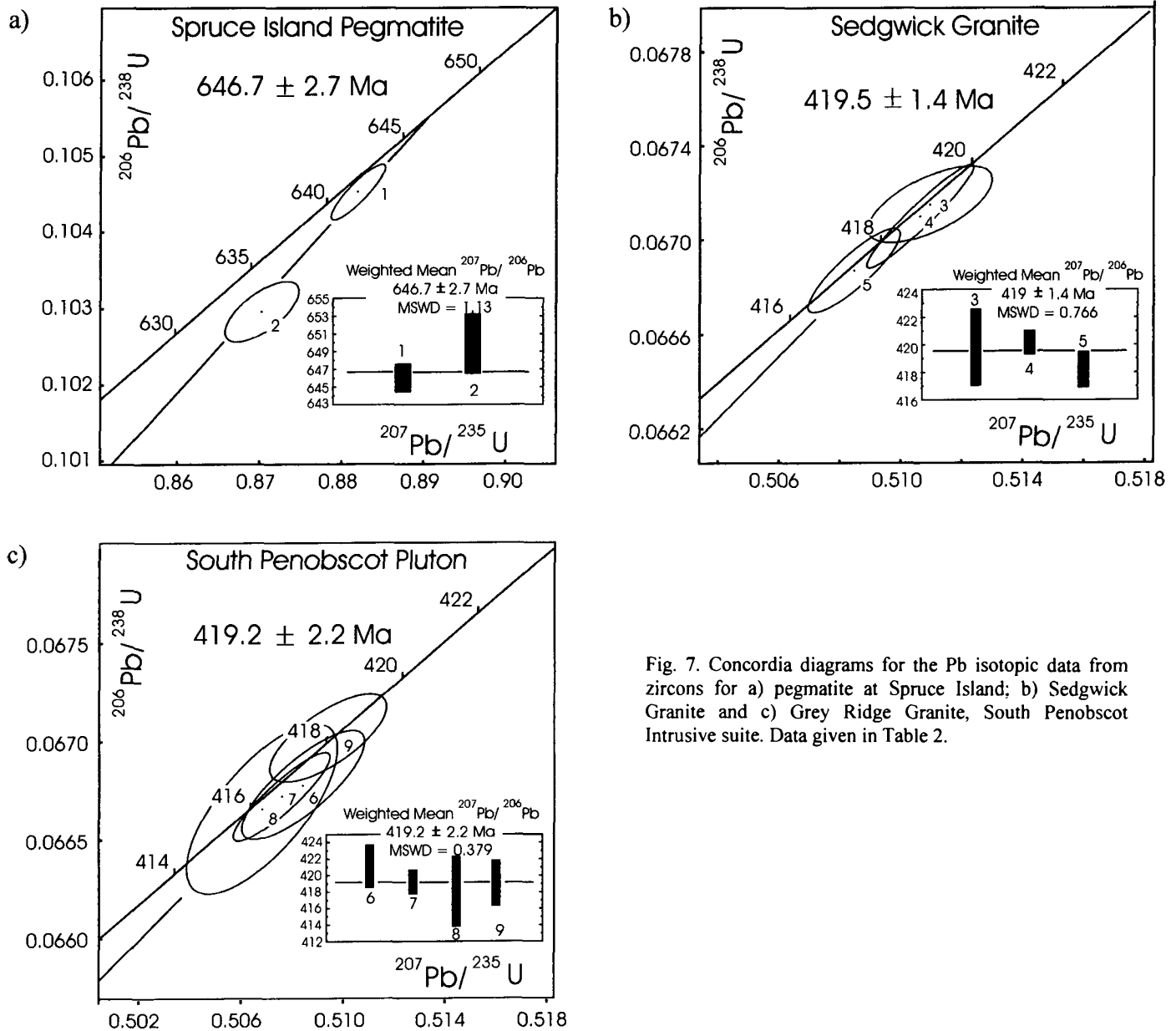


Fig. 7. Concordia diagrams for the Pb isotopic data from zircons for a) pegmatite at Spruce Island; b) Sedgwick Granite and c) Grey Ridge Granite, South Penobscot Intrusive suite. Data given in Table 2.

$^{40}\text{Ar}/^{39}\text{Ar}$ release spectra

Hornblende samples of very high purity weighing about 1 g and 0.025 cm to 0.0125 cm in size were separated from two amphibolite samples from the Seven Hundred Acre Island Formation that were collected at the west side of the south end of Spruce Island (Fig. 2). The suffix H added to the specimen number indicates that it is a hornblende separate. Sample S84-12 is foliated greenish black amphibolite with hornblende crystals up to 3 mm long. Sample S84-12A was collected about 10 m from S84-12 and is a greenish black garnet amphibolite with garnet grains up to 3 mm mostly pseudomorphed by chlorite.

The release spectra for both samples (Table 4 and Figs. 8a and 8b) are discordant and suggest they have suffered some, albeit minor, ^{40}Ar loss (Harrison and McDougall 1980). Some excess ^{40}Ar is also superimposed on the loss spectra, most

clearly shown in sample S84-12A-H (Fig. 8b). The uniformity of the K/Ca ratio (0.04 to 0.05) for the high-temperature steps indicates the high quality of the separates. The two highest temperature steps for sample S84-12-H yield an apparent plateau age of about $669 \text{ Ma} \pm 11 \text{ Ma}$, and the three highest temperature steps for sample S84-12A-H yield an apparent plateau age about $652 \pm 13 \text{ Ma}$. The two ages are indistinguishable within the uncertainty limits and the most straightforward interpretation is that these oldest ages in the loss spectra should approximate the time of cooling through the closure temperature after lower amphibolite facies metamorphism of the Seven Hundred Acre Island Formation. The loss-type spectrum suggests that a later episode of heating caused loss of ^{40}Ar (Harrison and McDougall 1980), possibly the later retrograde lower greenschist metamorphism.

Table 4. $^{40}\text{Ar}/^{39}\text{Ar}$ analyses for hornblendes from amphibolites (moles $^{39}\text{Ar} \times 10^{-14}$)

Temp °C	$\frac{^{40}\text{Ar}}{^{39}\text{Ar}}$	$\frac{^{37}\text{Ar}}{^{39}\text{Ar}}$	$\frac{^{36}\text{Ar}}{^{39}\text{Ar}}$	Moles ^{39}Ar	^{39}Ar % Tot.	% ^{40}Ar Rad.	K/Ca	Apparent Age (Ma)
S84-12-H; J value = 0.002882								
750	322.8	4.683	0.6961	2.1	1.6	36.4	0.1043	528 ± 23
800	272.6	4.654	0.5111	2.1	1.6	44.7	0.1049	545 ± 23
850	271.3	3.769	0.4370	2.7	2.0	52.5	0.1296	622 ± 19
900	293.4	3.303	0.5306	1.9	1.5	46.6	0.1480	601 ± 22
930	244.6	3.527	0.4199	3.9	2.9	49.4	0.1386	540 ± 10
975	334.3	9.093	0.7779	1.5	1.1	31.4	0.0535	480 ± 30
985	214.6	15.590	0.3054	4.8	3.6	58.5	0.0311	563 ± 15
1025	170.7	10.490	0.0937	14.4	10.8	84.3	0.0463	630 ± 6
1030	166.7	10.190	0.0708	20.1	15.1	87.9	0.0477	640 ± 6
1045	177.2	10.540	0.1088	11.4	8.6	82.3	0.0461	637 ± 7
1060	207.4	10.960	0.2099	24.6	18.5	70.5	0.0443	639 ± 6
1075	164.6	11.040	0.0349	28.5	21.5	94.3	0.0440	671 ± 6
Fuse	168.2	10.830	0.0517	14.7	11.1	91.4	0.0449	666 ± 6
Total gas age				132.8	100.0			636
Highest-temperature plateau age								669 ± 11
S84-12A-H; J value = 0.002924								
880	373.7	3.009	0.6499	3.3	1.9	48.7	0.1625	771 ± 14
930	286.4	2.206	0.5073	4.2	2.5	47.7	0.2218	607 ± 8
975	211.9	4.238	0.3277	6.0	3.5	54.4	0.1153	526 ± 12
990	226.1	6.768	0.3839	4.5	2.6	50.1	0.0720	518 ± 8
1010	206.3	9.324	0.2704	6.6	3.9	61.6	0.0522	574 ± 9
1025	170.2	10.680	0.1068	15.3	8.9	81.9	0.0455	621 ± 5
1040	177.7	11.090	0.1195	11.4	6.7	80.6	0.0438	636 ± 6
1050	179.5	11.260	0.1200	11.7	6.8	80.7	0.0431	642 ± 6
1060	181.2	11.670	0.1317	12.6	7.4	79.0	0.0416	636 ± 7
1080	177.0	11.780	0.1248	24.3	14.2	79.7	0.0412	628 ± 6
1100	166.2	11.890	0.0690	21.6	12.6	88.3	0.0408	649 ± 6
1125	162.6	11.870	0.0567	26.7	15.6	90.3	0.0409	649 ± 6
Fuse	163.9	11.760	0.0534	22.8	13.3	90.9	0.0413	657 ± 7
Total gas age				171.0	100.0			633
Highest-temperature plateau age								652 ± 13

TECTONIC SETTING OF THE ISLESBORO BLOCK

The data given here for the lithostratigraphy, geochemistry, geochronology, and metamorphic petrology for the Islesboro block augment earlier discussions of the tectonic history of the peri-Gondwanan terranes in the Penobscot Bay region by Stewart *et al.* (1995b), and by Stewart (1998).

Implied basement beneath the Seven Hundred Acre Island Formation is continental but of unknown age. The nature of continental crust related to the Islesboro block is unclear as neither Laurentian nor classical Avalonian basement *sensu stricto* is exposed in coastal Maine. The Pb isotopes of the Grenville massifs from Vermont (R. Ayuso and N. Ratcliffe, unpublished data) and from Ontario (Fig. 6) have lower values of $^{207}\text{Pb}/^{204}\text{Pb}$ (DeWolfe and Mezger 1994) than most of the more radiogenic rocks from the Islesboro block. The age-corrected isotopic differences for the estimated initial values (at 647 Ma) in the amphibolite and schist samples from the Seven Hundred Acre Island Formation and for Grenville gneisses show that the Grenville massifs cannot be the radiogenic source controlling the high $^{207}\text{Pb}/^{204}\text{Pb}$ values.

The nearest exposed Proterozoic basement is peri-Gondwanan and located in New Brunswick. In contrast to the Grenville massifs, the majority of basement rocks from peri-Gondwanan rocks and the Avalonian zone in the northern Appalachian orogen (Fig. 6) have radiogenic Pb (Ayuso and Bevier 1991; Ayuso *et al.* 1996; Ayuso and Arth 1997; Whalen *et al.* 1994a, 1994b, 1996a, 1996b; Ayuso 1999) such

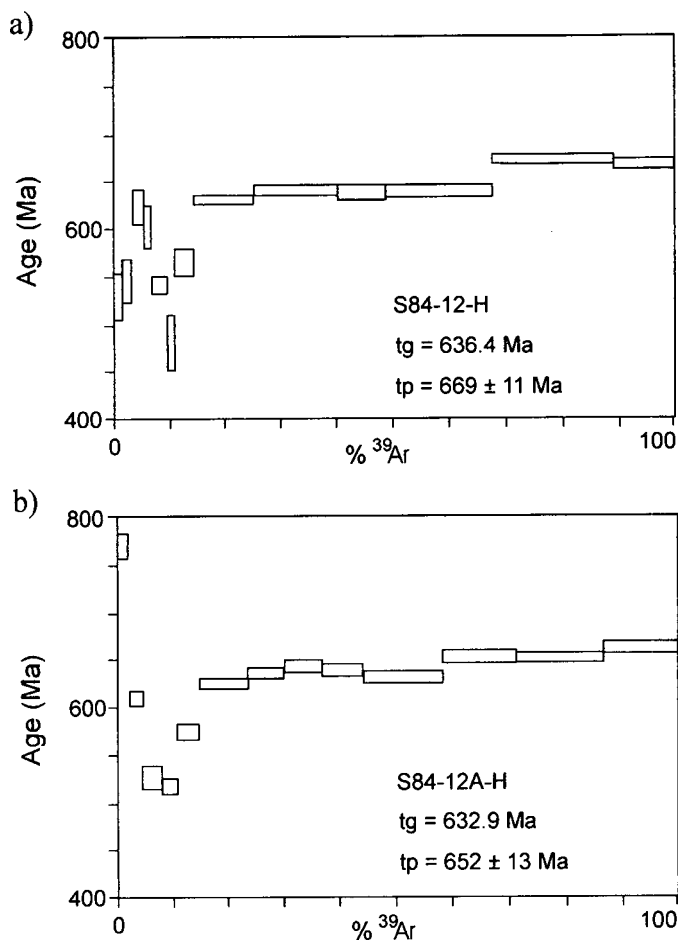


Fig. 8 $^{40}\text{Ar}/^{39}\text{Ar}$ release spectra for hornblende from amphibolite samples a) S84-12-H and b) S84-12A-H. Tg = total gas age; Tp = high temperature plateau age. Data given in Table 4.

as that required for the rocks from the Seven Hundred Acre Island Formation. Moreover, Paleozoic granitic rocks in this area are known to have incorporated evolved radiogenic crustal reservoirs, as suggested by ϵNd values as low as about -3.0 and T_{DM} values as old as about 2 Ga (Arth and Ayuso 1993; R. Ayuso unpublished data), and $^{207}\text{Pb}/^{204}\text{Pb} > \sim 15.6$ (Ayuso and Bevier 1991; Whalen *et al.* 1996a, b).

The protoliths of the Seven Hundred Acre Island Formation include mature quartz arenite, aluminous pelite, and dolomitic carbonate sediments that were deposited on a mature continental platform in the Late Proterozoic. The chondrite-normalized diagrams show that the schists are LREE-enriched and the HREE patterns are virtually flat. The Pb isotopic compositions of the whole rocks are more radiogenic than the average Pb crustal evolution curve (Stacey and Kramers 1975), and attest to their origin in evolved crustal isotopic reservoirs. The age of the Seven Hundred Acre Island Formation is probably Neoproterozoic, although not accurately known.

The continental crust was attenuated and the sediments upon it were intruded by sub-alkaline to alkaline basaltic rocks that were the protoliths of the amphibolite units studied. Immobile trace elements (Figs. 3, 4) and their ratios (Fig. 5)

allow inferences about the nature of the sources and tectonic settings. The amphibolite samples have flat REE patterns and show no evidence of fractionation of Eu by plagioclase or garnet. They are not related to arc sources as demonstrated by high abundances of Ti and Y, by low Th, and importantly, because they lack significant negative Nb (and Ta) anomalies with respect to Th and Ce (Pearce 1996). In general, the amphibolites have normalized patterns with negative slopes, Zr- and Ti-enrichment relative to MORB, and Y, Yb, and Sc values similar to MORB. These features are consistent with the involvement of MORB-type source regions. However, the high overall abundances of the REE and high field-strength elements compared to MORB and volcanic-arc basalts point to a predominant origin from an incompatible element-enriched (e.g., E-MORB) source or within-plate basalt-type source region (Fig. 5). The relative abundances of Y, Yb, and Sc are unlike those of typical within-plate basalts (typically depleted relative to MORB). Values of $Ti/Y > 100$ indicate that garnet was not in equilibrium with magmas represented by the amphibolite samples and attest to the contribution of deeper (>50 km) source regions as required for within-plate basalts. Thus, the amphibolite samples have geochemical variations and normalized patterns that are intermediate between tholeiitic and within-plate basalt end members consistent with derivation in a transitional setting where E-MORB magmas predominated and continental crust was under extension. The Pb isotopic compositions of the amphibolites are radiogenic and more similar to ratios observed in peri-Gondwanan basements in Atlantic Canada than in Grenville basement (Fig. 6).

The Seven Hundred Acre Island Formation was deformed and metamorphosed to lower amphibolite facies at about 670 to 650 Ma before it was intruded by the granitic pegmatite from which the zircon that was dated at 647 ± 2.7 Ma was obtained.

Late in the Proterozoic or early in the Cambrian the Islesboro Formation was deposited unconformably upon the Seven Hundred Acre Island Formation in an environment where a stable continental platform yielded clean quartz arenite, conglomerate, pelitic sand, and dolomitic limestone. There likely was a nearby volcanic source for volcanic-rich sediments in the Islesboro Formation. The Islesboro Formation was metamorphosed to lower greenschist facies, possibly at the same time that the underlying metamorphic rocks of the Seven Hundred Acre Island Formation were retrograded. This metamorphism may have been late in the Cambrian when the Ellsworth Schist was metamorphosed, but this hypothesis is speculative.

REGIONAL SIGNIFICANCE

The various peri-Gondwanan terranes were agglomerated into a composite terrane by a complex and incompletely known series of tectonic events between the Cambrian and the Silurian. This problem is especially difficult to decipher for coastal Maine and along strike to the northeast in New Brunswick (Stewart 1998; Fyffe *et al.* 1999, p. 77–81), and is being solved incrementally by new mapping (Johnson 2001),

discoveries of fossils (e.g. Fyffe and Riva 2001) and new results from Pb-zircon isotopic geochronology. The discussion that follows is focused on the region just described, but we have tried to make it consistent with what we understand about the tectonic history of southern New Brunswick.

Barr and White (1996) critically evaluated the tectonic and magmatic history of the many peri-Gondwanan as well as the Avalonian *sensu stricto* terranes in New Brunswick and Nova Scotia and concluded that peri-Gondwanan terranes there had basements up to ~200 Ma older than the Avalonian terranes *sensu stricto* and had different magmatic histories and different stratigraphic sequences in their lower Paleozoic cover sequences. They suggested that the suite of peri-Gondwanan terranes in Atlantic Canada comprised a tectonic agglomeration of terranes, possibly the equivalent of a Wilson cycle along the southeastern shore of the Iapetus Ocean and that may have acted as a buffer zone between the Avalon composite terrane and Laurentia during Silurian and Devonian continental collision. Cocks *et al.* (1997) discussed the agglomeration of different terranes along large continuous strike-slip motions during the late Proterozoic and early Paleozoic and tried to define as much of the history of these processes as they could using paleontological data for control. Johnson and McLeod (1996) describe highly faulted rocks in the New River belt of southern New Brunswick where Neoproterozoic granitoid rocks and volcanic rocks have lower Paleozoic cover sequences with some lithologic and chronological similarities to formations found in the Penobscot Bay region. Many papers like these describe the agglomeration and subsequent structural history of Avalonian and peri-Gondwanan composite terranes but we will consider only those aspects of the process we observe in northern Penobscot Bay, which seem to affect only peri-Gondwanan terranes.

The well-sorted and clean platform lithologies in the Islesboro block are similar to some rocks found in the lower part of the stratigraphic section in the St. Croix terrane (Fig. 1) exposed west of the Turtle Head fault zone. Berry and Osberg (2000) and Berry *et al.* (2000) have suggested that they may be correlative. However, unlike the lower greenschist rocks of the Islesboro Formation the rocks west of Penobscot Bay are metamorphosed to upper amphibolite facies and overlain by rusty black schist of Late Cambrian and Ordovician age that has not been found in the Islesboro block.

The oldest principal fault in the northern Penobscot Bay region (Fig. 1) is the Sennebec Pond thrust fault, which separates the peri-Gondwanan St. Croix terrane on the east from the peri-Gondwanan Fredericton trough terrane on the west. The age of folding and metamorphism of the St. Croix terrane must be either Ordovician or Early Silurian because the undeformed and unmetamorphosed Late Silurian Northport Granite and Youngtown Granite intrude metamorphosed rocks of the St. Croix terrane. Metamorphism of the rocks of the Fredericton trough terrane occurred before first motion on the Sennebec Pond fault because this fault separates lower and middle amphibolite facies rocks of the Fredericton trough west of Belfast. In places along the fault, metamorphosed rocks from both terranes are milled together and contain shear bands and asymmetric folds consistent with

dextral shear (Stewart *et al.* 1995, p. 45). The emplacement of the Sennebec Pond thrust fault and regional metamorphism of the adjacent terranes all may be of Silurian age. West *et al.* (1992) dated the Late Silurian (423 ± 3 Ma) Pocomoonshine pluton in eastern Maine and they interpreted it to intrude across a fault equivalent to the Sennebec Pond fault. This provides an upper constraint on the age of the Sennebec Pond fault.

After the emplacement of the St. Croix terrane, prominent younger strike-slip zones in the Silurian juxtaposed the peri-Gondwanan Ellsworth terrane and the smaller Islesboro block (Fig. 1). The older of these strike-slip zones, the inferred Penobscot Bay fault zone and its related splays, the Smith Cove and North Blue Hill faults, juxtaposed the Ellsworth terrane and the St. Croix terrane. In southern Islesboro, blocks of the mafic metavolcanic rocks of the North Haven Greenstone of the Ellsworth terrane were shuffled along strike-slip faults with blocks of the Islesboro terrane as the Islesboro block as a whole was moved northeastward between the Ellsworth and St. Croix terranes. About 10 km southeast of the present trace of the Penobscot Bay fault zone at Ames Knob in North Haven, Llandoveryan fossils of peri-Gondwanan affinity occur above an unconformity on North Haven Greenstone (Smith *et al.* 1907; Brookins *et al.* 1973; Berry and Boucot 1970; Gates 2001). The Penobscot Bay-Smith Cove fault cuts the contact metamorphic aureole around the 419.5 ± 1 Ma Sedgwick pluton (Fig. 7b). When movement on this fault zone was impeded by the emplacement of the South Penobscot Intrusive Suite the locus of right lateral strike-slip faulting shifted northward along other fault splays and isograds around the South Penobscot Intrusive Suite were displaced eastward. Continued strike-slip faulting was interrupted and the fault zone was re-established northwestward until it was replaced by the Turtle Head fault zone, which bounds the western side of the Islesboro block. The Turtle Head fault zone cuts the northern end of the South Penobscot Intrusive Suite, which was dated at 419.2 ± 2.2 Ma (Fig. 7c).

The Turtle Head fault strike-slip zone separates the greenschist facies rocks of the Islesboro block from the middle amphibolite facies regional metamorphic rocks of the St. Croix terrane, which were still above hornblende and biotite closure temperatures when fault motion occurred (West *et al.* 1995). However, no evidence of increased regional metamorphic grade has been observed in the Islesboro Formation east of the Turtle Head fault. The Turtle Head fault zone is intruded by the Lucerne Granite (380 ± 4 Ma, Zartman and Gallego 1979) and subsequent geologic events have not significantly changed the relationships. Later events include the attachment of the peri-Gondwanan and Avalonian composite terranes to Laurentia and several orogenic episodes in the Late Paleozoic that led to the formation of the Appalachian Mountains and in the Mesozoic that led to the formation of the Atlantic Ocean.

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APPENDIX A

Localities for samples (Fig. 2), and descriptions for samples listed in Table 1. Abbreviations are as follows: Ab, albite; Act, actinolite; And, andalusite; Bi, biotite; Ca, calcite; Ch, chlorite; Ctd, chloritoid; Ep, epidote; Ga, garnet; Hb, hornblende; Il, ilmenite; Mag, magnetite; Mu, muscovite; Or, orthoclase; Qu, quartz; St, staurolite; Ti, titanite; Zo, zoisite.

Amphibolite and hornblende gneiss

S84-3: Coarse amphibolite, southwestern coast of Lime Island. Mode: Hb 15%, Act 25%, Ch 25%, Ab 15%, Ep 10%, Qu 5%, Il 3%, Mag 2%.

Sample used for determination of whole rock Pb isotopes.

S84-12: Amphibolite, west side of Spruce Island at southwestern end.

Mode: Hb 70%, Ab 10%, Act 5%, Qu 5%, Ti 5%, Ch 2%, Mu 2%, Ap 1%.

Sample used for determination of whole rock Pb isotopes. Hornblende separate used for $^{40}\text{Ar}/^{39}\text{Ar}$ release spectra.

S84-12A: Garnet amphibolite, 10 m from site of S84-12, west side of Spruce island at southwestern end.

Mode: Hb 70%, Qu 10%, Ch 10%, Ga 5%, Ti 3%, Ep 2%, traces of Ap and Il.

Sample used for determination of whole rock Pb isotopes. Hornblende separate used for $^{40}\text{Ar}/^{39}\text{Ar}$ release spectra.

S84-13: Amphibolite, northernmost point of Spruce Island.

Mode: Hb 35%, Act 10%, Ab 40%, Ep 5%, Ch 5%, Ti 3%, Mag 2%.

S84-19: Amphibolite from deformed dike in marble, southwestern point on Lime Island.

Mode: Hb 25%, Act 30%, Ab 25%, Zo 15%, Ch 3%, Mag 2%. Sample used for determination of whole rock Pb isotopes.

S68-175: Hornblende-rich gneiss, middle of east side of Spruce Island.

Mode: Hb 90%, Ep 8%, Ch 2%, traces of Mag.

Sample used for determination of whole rock Pb isotopes.

Schist and gneiss

S68-65: Garnet-muscovite-andalusite-staurolite schist, west of boat landing, Cradle Cove, Seven Hundred Acre Island.

Mode: Ga 30%, Mu 30%, Qu 30%, And 5%, St 2%, Ctd 2%, traces of zircon, rutile and pyrrhotite.

Sample used for determination of whole rock Pb isotopes.

S68-177: Splendid garnet gneiss, southern end of Spruce Island.

Mode: Mu 62%, Qu 25%, Ga 5%, Ch 2%, And 2%, Mag 1%.

Sample used for determination of whole rock Pb isotopes.

S84-23: Splendid muscovite schist, small cove off east side of Seal Cove, west of road.

Mode: Mu 25%, Qu 25%, Ga 20% (mostly replaced by Ch, with layers of quartz grains), sericite groundmass 25%, Mag 5%.

Sample used for determination of whole rock Pb isotopes.

S84-17: Siliceous carbonate gneiss with pseudomorphs after idocrase(?), west side of the north point, Spruce Island.

Mode: Ca 55%, Bi 20%, Idocrase (pseudomorphed to sericite) 15%, Qu 7%, Mag 3%.

Muscovite pegmatite

S84-18: Muscovite pegmatite, intrudes marble, eastern part of north point on Spruce Island.

Mode: Or 35%, Ab 25%, Qu 23%, Mu 7%.

Zircon from this sample was used for U-Pb age dating.